Aquifer Vulnerability Assessment at Ipinsa-Okeodu Area, Near Akure, Southwestern Nigeria, using GODT

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Abstract

Aquifer vulnerability assessment has been undertaken at Ipinsa and Oke-Odu area, southwestern Nigeria using geoelectrically derived GODT model. One hundred and two (102) vertical electrical soundings (VES) data utilizing Schlumberger array were carried out at half-current electrode separation (AB/2) varied between 1-150 m. Qualitative interpretation of the VES data using partial curve matching and computer-aided inversion techniques yielded geo-electric parameters (layer resistivity and thickness) that were used in delineating the aquifers in the area and evaluating their vulnerability to pollutants. The geoelectric sections revealed that the area is underlain by three to four geo-electric layers namely the topsoil, weathered layer, partly weathered/fractured basement and fresh basement. The weathered layer and partly weathered/fractured basement constitute the major aquifers in the area. The GODT vulnerability model depicts that the area is characterized by four vulnerability zones which are very low, low, moderate and high vulnerable zones. According to the model, about 10% of the area is highly vulnerable while about 35% is of moderate rating. The low and very low ratings constitute 40% and 15% of the area respectively.

Keywords: Aquifer vulnerability, GODT model, vertical electrical sounding and geoelectric parameters.

1. Introduction

The search for groundwater has been on the increase across the globe and this is because it is considered all over the world to be the best source of potable and safe both for drinking, agricultural and industrial purposes (Hoque et al., 2009). Several tools ranging from geophysical, remote sensing, geographic information system (GIS) among others, have been utilized for proper location of this precious resource within the subsurface. However, location of this resource is not enough without proper protection from harmful substances that can degrade its quality and render it unfit for human use. Groundwater reservoirs are easily affected by pollution through a process that is slow and consequently dreadful (Baghvand et al., 2010). Prevention of aquifers pollution is considered as an important factor in the management of groundwater resources and as such aquifer vulnerability study becomes imperative within the domain of groundwater study. Hence, the assessment of groundwater vulnerability to pollution has been the subject of intensive research during the past years and a variety of index methods have been developed to evaluate aquifer vulnerability. These methods include DRASTIC (Aller et al., 1987), GOD (Foster, 1987), AVI (Van Stempvoort et al., 1993), SINTACS (Civita, 1994) e.t.c. and are all subjective to varied vulnerability parameters. The derivation of the various parameters required for the computation of the index vulnerability models is usually multi-disciplinary while the accuracy of the resulting models depends majorly on the available information and their authenticity. Meanwhile, site specific vulnerability assessments using these methods are not readily feasible since in most cases there might not be enough hydrogeological information to compute and thus they are usually applied at regional scale. Consequently, attempt is made in this study to compute hydrogeological parameters from geoelectric parameters for the assessment of aquifer vulnerability at Ipinsa and Oke-Odu area, near Akure, Southwestern Nigeria. Most geophysical assessments of aquifer vulnerability recorded in the literatures have engaged the use of longitudinal conductance, a second order geoelectric parameter to assess the protective capacity of the overburden units, (Abiola et al., 2007, Aweto, 2011; Akintorinwa and Olowolafe, 2013). This approach however, is insensitive to the possible presence of relatively high resistive geological formations like laterites that are good protective barriers for the underlying aquifers. More importantly, vulnerability models become more effective as more parameters influencing the disposition of contaminants are available as input for the model. Therefore, four aquifer vulnerability parameters namely groundwater occurrence (G), overlying strata (O), depth to aquifer (D) and topography (T) are integrated in this study to assess the aquifer vulnerability of the study area. The former three parameters GOD has been successfully integrated for aquifer vulnerability assessment in the past (Foster, 1987; Khemiri et al, 2013) while the fourth parameter (T, topography) is an added input parameter considered to improved the resulting vulnerability model since the topography of an area can influence the migration of contaminants. The ridges usually associated with run-off and less infiltration, while the opposite is the case for depression. Furthermore, studies have shown that contaminants can be topographically controlled whereby contaminants are held downslope by gravity and prevented from migrating upslope (Khemiri et al., 2013).

2. Site Description

The study area covers two communities namely; Ipinsa and Okeodu situated near Akure Ondo State (Figure 1). It

lies within latitudes 7⁰ 17⁰ 44.7¹N and 7⁰ 19⁰ 21.9¹N and longitudes 5⁰ 07⁰ 49.23¹E and 5⁰ 09⁰ 37.25¹E. The study area occupies a total area of about 10 km². The terrain across the study area is undulating with surface elevation ranging between 355 m and 430 m above sea level with more depressions in the southeastern part relative to the northwestern part (Figure 2). The study area is underlain by the Precambrian Basement Complex rocks of Southwestern Nigeria. The two lithologic units recognized in the area include; undifferentiated Older Granite-Charnockites suites and Migmatite-Gneiss-Quartzite complex (Figure 3). The undifferentiated Older Granite-Charnockites suites are located in the southeastern part of the area and occur as extensive low-lying outcrop while the Migmatite-Gneiss-Quartzite complex essentially occupies the northwestern part of the area (Figure 3).

3. Methodology

One hundred and two (102) Vertical Electrical Soundings (Figure 1) data were acquired in the study area, using PASI 16GL Earth Resistivity Meter and its accessories. The Schlumberger array was adopted for the field survey, with half current electrode spacing (AB/2) varying from minimum of 1 to maximum of 40 to 150m depending on the depth to bedrock and spread allowance. Geoelectric sounding data were interpreted manually using the conventional partial curve matching technique involving the use of theoretical and auxiliary curves (Keller and Frishchnecht, 1966; Koefoed, 1979). The derived geoelectric parameters were further refined using a forward modelling computer algorithm, WinRESIST Version 1.0 (Vander Velpen, 2004). The geoelectric results were presented as curve types, distribution charts, and maps.

The GODT index which is used to evaluate the aquifer vulnerability in the area was calculated by multiplication of the influence of the four parameters such as Groundwater occurrence (confinement of the aquifer), Overall lithology overlying the aquifer, Depth to the aquifer and Topography of the area. These GOD parameters were interpreted from the geoelectric parameters (resistivity and thickness of the interpreted layers) while the topography was obtained from the surface elevations recorded in the area. Values from 0 to 1 were as assigned to these parameters as shown in Table 1.



Figure 1: Base map of the study area showing the Vertical Electrical Sounding (VES) Stations Outset Administrative map of Nigeria, (After Obaje, 2009)

Tuble 1. That builds for GOD T mouel parameters (mouthed after Themint et al., 2015)							
Aquifer Type	Note	Depth to Aquifer (m)	Note	Lithology (Ω-m)	Note	Topography	Note
Non-Aquifer	0	<2	1	<60	0.4	Ridge	0.7-0.8
Artesian	0.1	2-5	0.9	60-100	0.5	Depression	0.9-1
Confined	0.2	5-10	0.8	100-300	0.7		
Semi-confined	0.3-0.5	10-20	0.7	300-600	0.8		
Unconfined	0.6-1	20-50	0.6	>600	0.6		
		50-100	0.5				

Table 1. Attribution of notes for GODT model parameters (modified after Khemiri et al., 2013)

The GODT index was then calculated by multiplying the influence of the various parameters together as

1

shown in equation 1 GODT Index = $G \times O \times D \times T$ Where: G = Type of Aquifer O = Overburden Lithology D = Depth to the Aquifer T= Topography





4. Results and Discussion

The summary of the interpreted results of the VES curves at each VES stations are as presented in Table 2. The characteristic curve types obtained in the area are A, H, K, HA, AA, QH, HK, AK, QKH, AKH and HKH. Figure 4 showed the order of predominance of the curve types obtained in the study area. The geoelectric sections generated along north-south and west-east directions revealed that the study area is underlain with four subsurface layers corresponding to topsoil, weathered layer, partly weathered/fractured basement and fresh basement with resistivity values in the range of (37-257 ohm-m), $(33-784 \Omega \text{m})$, $(710-772 \Omega \text{m})$ and $(1614-9025 \Omega \text{m})$ respectively (Figures 5 and 6). The weathered layer and the partly weathered/fractured basement constitute the major aquifers in the study area. The topsoil on the other hand, constitute majorly the protective cover for these underlying aquifer but is generally thin (0.9-2.7 m) and characterized by relatively high resistivity values within the range of 60-600 Ω m which is considered to offer poor to moderate protection for the underlying aquifers. However, it could be observed that about 30% of the study area is characterized by low resistivity values (< 60 Ω m) and high resistivity values (> 600 Ω m) which suggest good protective capacity of the underlying aquifers (Figure 7).

Figure 8 shows the GODT vulnerability model generated based on four parameters: i) G, groundwater confinement, ii) O, overlying strata, iii) D, depth to the aquifer and iv) T, topography of the area. The ultimate integrated aquifer vulnerability index is the final product of component indices for these parameters (Foster *et al.*, 2002; Afonso *et al.*, 2008). The study area was categorized into four ratings viz; very poor, poor, moderate and

high vulnerability zones based on the Vulnerability assessment presented in Table 4.3 (Murat et al., 2003). The vulnerability model shows that major part of the study area falls within the low and moderate vulnerability classes. The most vulnerable zones transect the southeastern axis of the area where low surface elevations are recorded. Thus, the aquifers in these areas are adjudged readily vulnerable to contamination from near surface pollutants.

5. Conclusion

Aquifer protection is essential for a sustainable use of the groundwater resources, protection of the dependent ecosystems, and a central part of spatial planning and action plans. The key expression for a quantification of aquifer protection is vulnerability. It is in view of this that this research has been undertaken to effectively characterize the vulnerability of the ambient aquifers to near surface contaminants around areas connecting Ipinsa-Okeodu, near Akure, Southwestern Nigeria. The GODT vulnerability model depicts that the study area is characterized by four vulnerability zones which are very low, low, moderate and high vulnerable zones. According to the model, about 10% of the area is highly vulnerable while about 35% is of moderate rating. The low and very low ratings constitute 40% and 15% respectively of the area. Therefore, it is highly recommended that the least vulnerable zone should be the primary target for future groundwater development in the area in order to ensure continuous supply of safe and potable groundwater for human consumption in the area and more importantly, location of septic tanks, petroleum storage tanks, shallow subsurface piping utilities and other contaminant facilities should be confined to these least vulnerable zones.

VES No	Layer Thickness (m)			Layer R	Curve					
	h_1	h ₂	h ₃	h4	ρι	ρ2	ρ3	ρ4	ρ5	Туре
1	2.7	7.3			238	*58	710			Н
2	1.1	0.4	5.5		371	4135	*140	2769		KH
3	1.3	3.8	15.5		237	586	*148	561		KH
4	0.9	3.3	2.6	34.7	263	136	583	*162	x	НКН
5	0.7	4.2	30.0		761	704	*113	4677		QH
6	0.7	1.0	3.9	26.7	348	110	1843	*96	4184	HKH
7	1.0	8.4	6.8		1142	1474	*570	1403		KH
8	1.0	3.1			434	*53	772			Н
9	0.9	8.9			391	*187	9025			Н
10	4.0	14.8			1889	851	*108			Q
11	1.0	3.0	11.3		213	448	*255	1940		KH
12	1.0	4.0	9.1	18.3	171	332	1531	*228	1432	AKH
13	1.3	0.6	9.0	14.7	437	158	721	*175	712	HKH
14	0.6	2.6	18.2		40	168	*70	708		KH
15	0.8	3.0			234	717	*82			K
16	1.0	1.7	13.1		143	1685	*30	1252		KH
17	1.0	10.8			101	573	*96			K
18	0.8	6.5	28.5		274	784	*33	98		KH
19	0.8	2.9	25.8		143	682	*131	820		KH
20	0.9	8.9	27.0		77	224	*128	1010		KH
21	2.5	4.5	11.0		483	152	*52	1039		QH
22	8.4	9.5			133	*74	3614			H
23	4.5	14.5	52.3		474	208	*46	x		QH
24	9.9	1.9	13.6		692	1996	*63	5333		KH
25	1.1	12.6			159	767	*181			K
26	1.6	12.4			154	971	∞			А
27	4.6	7.9			73	*108	523			А
28	1.7	2.9	16.6		113	661	*60	1027		KH
29	0.6	8.2	5.0		82	*264	*58	638		KH
30	0.8	4.7	20.0		84	420	*86	293		KH
31	1.0	2.0	5.3		240	840	*416	3080		KH
32	1.0	1.7			47	30	*398			Н
33	0.9	5.1			67	*32	380			Н
34	0.8	1.9	22.9		233	401	*92	785		KH
35	0.9	3.3			122	*15	232			Н
36	0.6	0.8	9.3		37	167	*107	2495		KH
37	1.2	2.8			143	*91	2859			Н
38	2.0	11.2			121	*77	270			Н
39	1.1	5.0	1	1	92	*19	587			Н
40	0.7	0.4	9.8		178	117	*197	69		HK
41	1.6	1.8	4.3		80	195	*21	177		KH
42	1.1	4.0	11.2		103	491	*43	305		KH
43	0.6	6.5	28.0		86	377	75	817		KH

Table 2. Summary of VES results

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44	1.2	5.7			150	*211	8853			Н
45	2.8	9.8			506	*188	4323			Н
46	1.0	2.4			283	*183	1829			Н
47	1.6	0.4			93	301	2635			А
48	1.0	6.9	29.4		198	579	*111	285		KH
49	1.8	5.0	19.0		150	479	1831	287		AK
50	1.0	2.5			72	3034	4166			А
51	0.8	2.8			106	592	1782			А
52	1.2	31.8			98	3163	1205			К
53	1.7	11.3	24.4		212	350	*333	1125		КН
54	0.9	1.8			128	181	2352			A
55	03	14.4			170	1148	*313			K
56	1.0	87			237	1177	*171			K
57	0.5	49	28.3		161	246	*85	1614		КН
58	0.9	2.1	20.5		810	1582	555	1014		KII
50	0.8	2.1 6.4	83		72	264	*07	230		КН
59	0.5	2.7	0.5		286	*161	5152	239		
61	1.0	3.7			457	160	2607			П Ц
62	1.9	1.9	2.2	10.5	437	100	2097	*17	260	П
62	0.0	0.9	2.3	10.5	227	43	/89	*4/	309	НКП
03	1.5	/./	28.2		176	898	*200	038		KH VU
64	0.9	4.1	28.2		1/0	413	*69	934		KH
65	0.7	3.0	27.1	10.1	190	300	*86	988	720	KH
66	0.8	3.0	4.8	19.1	16/3	492	6658	*38	/39	HKH
6/	1.0	3.5	4/.6		99	353	*199	1403		KH
68	0.7	10.2	26.5		193	528	*/2	318		KH
69	1.5	4.8	25.2		54	314	*68	10/8		KH
70	0.7	2.0	13.2		/10	803	*148	515		KH
/1	0.8	3.1	20.2		8/	*62	3064	1010		H
72	0.6	4.2	29.2		56	88	*281	1018		AA
/3	1.0	6.5			/0	*132	1310			A
/4	0.8	1.5			129	390	2103			A
75	0.7	30.1			40	104	1070			A
70	0.5	2.2			40	41	1070			A
70	1.0	5.5	12.1		40	(72)	*25	020		
78	0.5	0.2	13.1		250	0/2	*33	930		КП
79 80	2.0	30.3	20.0		239	2409	300	1002		
80	0.2	11.5	30.0		20	*296	1499	1005		А
01 92	0.5	3.0	2.5	12.0	566	200	1400	*470	1840	A
02	0.0	1.9	3.5	13.0	26	200	2262	-4/9	1049	
83	0.7	1.3			154	30 *190	2202			A
85	0.2	13.5	0.7		167	228	*160	2070		A VU
86	1.0	4.2	2.1		70	*87	100	2070		Λ
87	0.8	1.2			69	1020	1896			Δ
89	0.8	2.0			25	1020	1890			Λ Λ
80	2.2	2.0			146	4//· 600	1744			A
00	1.0	4.3			77	216*	1/44			A
90	1.2	4.2			17	210*	1110 ~			A
02	1.0	4.2	0.5		42	504	121*	~		A VU
92	1.0	0.5	9.5		121	75*	2200	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~		
93	0.0	4.6	74		121	31	87*	~		НΔ
95	17	9.0	/		80	32/1*	2645	~~		
95	3.8	7.7 1.8			86	145	2045			Δ
97	0.6	0.6	4.0		140	5603	~ 484*	25334		KH
98	0.0	1.2	8.2		79	335	101*	20004		КН
99	0.0	7.2	0.2		38	208*	191 20	~~		Δ
100	0.8	7.5			211	104*	~ ~			Н
101	0.0	1.5	3.9		76	122	∞ 94*	~		КН
102	2.5	49	5.9		87	166*	5942	~		IXII

*Major Aquifer



Figure 4: Frequency Distribution of Curve Types Obtained in the Study Area



Table 3	: Interval	Values	of the	GODT	Index a	nd Corres	ponding	Classes	(Modified	after	Murat	et al
2003).												

Index	Vulnerability Class
0-0.1	Very Low
0.1-0.3	Low
0.3-0.5	Moderate
0.5-0.7	High
0.7-1.0	Very High





References

- Abiola O., Enikanselu P. A. and Oladapo M. I. (2009). Groundwater Potential and Aquifer Protective Capacity of Overburden Units in Ado-Ekiti, Southwestern Nigeria; International Journal of Physical Sciences, 4 (3) 120-132.
- Adagunodo, T. A. and Sunmonu L. A. (2012). Geoelectric Assessment of Groundwater Prospect and Vulnerability of Overburden Aquifers at Adumasun Area, Oniye, Southwestern Nigeria. Arch. Appl. Sci. Res., 4 (5):2077-209.
- Afonso, M. J., Pires, A., Chamine, H. I., Marques, J. M., Guimares, L., Guilhermino, L. and Rocha, F. T. (2008). Aquifer Vulnerability Assessment of Urban Areas Using A GIS-Based Cartography: Paranhos Groiundwater Pilot Site, Porto, NW, Portugal. 33rd Int. Geological Symposium: Hydrogeology, Oslo (Norway).
- Aller, L., Bennet, T., Lehr, J.H., Petty, R.J. and Hackett, G. (1987). DRASTIC: A Standard System for Evaluating Groundwater Pollution Potential using Hydrogeologic Settings. EPA/600/2-85/018, US Environmental Protection Agency, Ada, Oklahoma, 455pp.
- Akintorinwa O. J. and Olowolafe, T. S. (2013). Geoelectric Evaluation of Groundwater Prospect within Zion Estate, Akure, Southwest, Nigeria. International Journal of Water Resources and Environmental Engineering. Vol. 5(1). Pp. 12-28

- Aweto, K. E. (2011). Aquifer Vulnerability Assessment at Oke-Ila area, Southwestern Nigeria. International Journal of the Physical Sciences Vol. 6(33), pp. 7574 7583,
- Baghvand, A., Nasrabadi, T., Nabibidhendi, G., Vosoogh, A., Karbassi, A., Mehradadi N (2010). Groundwater Quality Degradation of an Aquifer in Iran central desert. Desalination 260(3):264-275.
- Civita, M., 1994. Le Carte della Vulnerabilità degli acquiferi all inquinamento: Teoria and pratica. Pitagora Editrice, Bologna
- Foster, S.S.D. (1987). Fundamental Concepts in Aquifer Vulnerability Pollution Risk and Protection Strategy. In Vulnerability of soil and groundwater to pollution: Proceedings and information. W. van Duijvenboodennd H.G. van Waegeningh (editors).TNO Committee on Hydrological Research, The Hague, 69-86.
- Harter, T. (2003). Groundwater Quality and Groundwater Pollution. Publication 8084, http://anrcatalog.ucdavis.edu.
- Hoque, M. A., Khan, A. A., Shamsudduha, M., Hossain, M. S., Islam, T. and Chowdhury, S. H. (2009). Near Surface Lithology and Spatial Variation of Arsenic in the Shallow Groundwater: Southeastern Banglandesh. Environmental Geology, 56, 1687-1695.
- Keller, G. V. and F. C. Frishchnecht, (1966). Electrical Methods in Geophysical Prospecting. Pergamon Press, New York, pp. 96.
- Koefoed, O., (1979). Geosounding Principles 1. Resistivity Measurements. *Elsevier Scientific Publishing, Amsterdam, Netherlands.*, pp. 275.
- Murat, V., Paradis, D., Savard, M.M., Nastev, M., Bourque, E., Hamel, A., Lefebvre, R. and Martel, R., (2003). Vulnérabilité à la nappe des aquifères fractures du sud-ouest du Québec- Évaluation par les methods DRASTIC et GOD. Current Research, no. 2003-D3, 2003; 14 p.
- Obaje, N.G. (2009). Geology and Mineral Resources of Nigeria. Published by Springer London. 221p.
- Omosuyi, G. O. (2010). Geoelectric Assessment of Groundwater Prospect and Vulnerability of Overburden Aquifers at Idanre, Southwestern Nigeria. Ozean Jour. App. Sci. 3(1). Pg.19-28.
- Omosuyi, G.O.and Oseghale, A. (2012). Groundwater Vulnerability Assessment in Shallow Aquifers using Geoelectric and Hydrogeologic Parameters at Odigbo, Southwestern Nigeria. Am. J. Sci. Ind. Res., 3(6): 501-512
- Vander Velpen, B. P. A., (2004). WinRESIST Software Version 1.0. ITC, IT-RSG/GSD, Delft, Netherlands.
- Van Stempvoort D, Ewert L, Wassenaar L (1993). Aquifer Vulnerability Index (AVI): A GIS Compatible Method for Groundwater Vulnerability Mapping. Can Water Res J 18:25–37.